

Author	What?	Where?	Results	Details of results
Lu et al. 2009	Aerosol-cloud-drizzle relationship	Monterey California (MASE-II experiment)	<ol style="list-style-type: none"> 1) Increase of N_c and decrease of r_e are associated with increase subcloud N_a. 2) $N_c \propto 1/d$ 3) $R_{cb} \propto H^3/N_c$ 4) Cloud precipitation susceptibility to aerosol perturbation $So' = 0.60-0.63$ (satellite derived value) 	<ol style="list-style-type: none"> 1) Relation is enhanced in close to adiabatic LWC cases (constant LWC: Twomey effect) 2) Relation is enhanced in non-adiabatic conditions. Greater d near the cloud base due to downdraft entrainment and/or depletion of water by drizzle. 3) Corroboration of the inverse relation with observational data. 4) $So' = -d \ln R / d \ln \alpha$ (at fixed LWC)
Comstock et al. 2004	Reflectivity and rain rate as a function of drizzle, N_c , and mean size	Southeast Pacific Ocean: EPIC experiment	<ol style="list-style-type: none"> 1) $Z = aR^b$ where b is related to the relative variability of N_c and r 2) Drizzle rate in Sc are dependent on both LWC and N_c 3) Inverse relationship between R and N_c 	<ol style="list-style-type: none"> 1) a and b values are inferred from different independent datasets.
Khairoutdinov et al. 2000	<ol style="list-style-type: none"> 1) N_c, cloud/drizzle mixing ratios, cloud drizzle N_c, cloud drop integral ratios. 2) Cloud water sources and sinks by drizzle are regressed using the N_c size spectra 	Large Eddy Simulation	<ol style="list-style-type: none"> 1) In drizzle conditions N_c and N_a decreases. 2) Drizzle initiation depends on the prediction of N_c and LWC. 3) Drizzle fallout rate (drizzle dwelling time in cloud) is a function of drizzle drop concentration. 	<ol style="list-style-type: none"> 1) Introduction of N_a as a variable from which N_c is predicted
Pawlowska and Brenguier 2002	Aerosol-cloud-drizzle relationship (in a suitable scale for GCM)	Eastern Atlantic Ocean: ACE-2 experiment	<ol style="list-style-type: none"> 1) Auto-conversion threshold 2) Existing parameterization schemes 3) Grid scale diagnostic of the precipitation rate 	<ol style="list-style-type: none"> 1) The two pristine cases, exhibit N_d reduction by droplet collection, while the three polluted cases exhibit smaller drizzle concentrations and predominantly in association with comparable or larger N_d values than N_{act}. 2) It appears that the drizzle reduction rate by precipitation is accurately represented by a H^3/N_{act} 3) Precipitation process in extended boundary layer clouds could be treated in GCM with a diagnostic scheme based on only H and N_{act}.
Lu et al. 2007	<ol style="list-style-type: none"> 1) Ship track effect on N_c and r: direct effect 2) Ship track effect N_d and LWC: indirect effect 	Monterey California: MASE-II experiment	<ol style="list-style-type: none"> 1) The ship track region exhibited a smaller r_e, reduced N_d, and larger cloud LWC than the adjacent clean region. 2) As subcloud N_a increase N_d increases. 3) Smaller and more numerous drizzle drops were found near the cloud top. 4) The smaller cloud base drizzle rate does not result in a larger LWP when compared with the clean clouds. 	<ol style="list-style-type: none"> 1) Evidence for the first indirect effect, and indirect effect. 2) Robust evidence of Twomey effect 3) The data suggest that more polluted cloud have fewer embryonic drizzle drops near cloud top, resulting in a smaller cloud base drizzle rate. 4) It appears that direct conversion of cloud drops to drizzle is insufficient to explain the dependence of LWP on aerosol number concentration.

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Stevens et al. 2005	Pockets of open cells and drizzle	1) Southeast Pacific: EPIC experiment 2) Northeast Pacific: DYCOMS-II experiment.	Long-lived regions of open cellular convection are coupled to the development of precipitation	1) Observation (satellite imagery) of POC formation from a region of open cellular convection. 2) Values near zero of ΔT_B within the POC region. 3) the largest precipitation rate (1 cm day^{-1}), in DYCOMS-II observations, occurs within the POC 4) Precipitation from Sc drives a more diabatic (cumulus like) circulation with narrower regions of convective ascent
Agee 1984	Satellite observation to study dynamic climatology: Mesoscale cellular convection		1) MCC occurs primary to the east of continents over warm ocean or to the west of continent over cold ocean. 2) Two convective mechanisms of MCC: heating from below and cooling from above	1) Cases of study in Steven et al. (2005): Southeast Pacific and Northeast Pacific, correspond to the latter case. 2) The the west of continent MCC formation respond to the cloud top cooling and inferred destabilization
van Zanten et al. 2004	Nc-drizzle relationship Drizzle rate variation with brightness temperature and re	Northeastern Pacific: DYCOMS-II experiment	1) The averaged drizzle rates are mainly due to the occurrence of localized patches of strongly enhanced precipitation. 2) Variability of drizzle rate correlate well with $R_{cb} \propto H^3/N_c$. 3) On the macroscopic scale the drizzle rate also correlates negatively with T and shows signs as well of thresholdlike dependence on in situ cloud-top effective radius.	1) This strongly suggests that drizzle could induce a transition in cloud structure 2) Thus higher precipitation rates are not due to a change in shape of the drizzle drop distribution but are mainly caused by the more frequent occurrence of larger drizzle drops. 3) $\Delta T = \text{difference}$ between the 11-m and 14-m brightness temperatures as measured by the <i>GOES-10 satellite</i>
Bretherton et al. 2004	Nc-drizzle relationship	South East Pacific Ocean: EPIC experiment	1) Less drizzle falling fro a cloud of a given thickness during period of high droplet number concentrations	1) Results are derive from LWC estimated from microwave radiometer and a estimate of the cloud droplet concentrations.
Stevens et al. 2003		Northeast pacific: DYCOMS-II experiment	1) CTEI=0.45 Value significantly larger than the critical value of 0.23 thought to portend the break-up of the cloud layer. 2) entrainment rate of 0.39 com s^{-1} . 3)LES was evaluated against the observational data, reproducing in good agreement the macroscopic evolution of the cloud layer.	1) CTEI stability parameter is derived from comparison of the STBL and Free-tropospheric states . 2) From consistent quasi-independent estimates. 3) Difficulty in maintaining the observed inversion-layer structure). Further work is necessary using LES with more accurate representations of long-wave radiative transfer